Comment on “Geochemistry of buried river sediments from Ghaggar Plains, NW India: Multi-proxy records of variations in provenance, paleoclimate, and paleovegetation patterns in the late quaternary” by Ajit Singh, Debajyoti Paul, Rajiv Sinha, Kristina J. Thomsen, Sanjeev Gupta

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Singh et al. (2016) published a geochemical record of sediment compositions from the flood plain of the Ghaggar River in western India and use the changing provenance, particularly as traced by Nd isotope composition, to reconstruct how erosion patterns have changed over the past 100 k.y. In doing so they propose a link between climate change and erosion, and they argue for more erosion from the Higher Himalaya during warmer interglacial periods and more from the Lesser Himalaya during glacial intervals. While we support the concept of erosion patterns being climatically modulated we here take the opportunity to culminating in Pakistan) near Marot we see that they form as single progressive trend, of the Hakkra River (downstream equivalent of the Ghaggar in modern culminating in Pakistan) near Marot we see that they form as single progressive trend, of the Hakkra River (downstream equivalent of the Ghaggar in modern

Sinha et al. (2016) have been consistently more positive than known modern Sutlej compositions since ~17 ka (~17 to ~19) (Alizai et al., 2011; Tripathi et al., 2004) also indicates that the drift to more positive values prior to 17 ka could be Himalayan (Sutlej) derived, but such flow is known to be finished by 5.2−5.7 ka at the latest (Clift et al., 2012). The fact that the εNd values of the sands considered by Singh et al. (2016) were consistently more positive than known modern Sutlej compositions since ~17 ka (~17 to ~19) (Alizai et al., 2011; Tripathi et al., 2004) also indicates that the drift to more positive values was not caused by changes in the Sutlej basin, which is now more negative, but to mixing with another sediment source. The trend
in Ghaggar-Hakkra εNd values is not directed towards a known modern river that could actually have received material from the Lesser and Higher Himalaya. Instead the trend since 14 ka has been away from the Himalayan draining tributaries but towards the average of the Thar Desert (Fig. 1), similar to the modern Ghaggar, which is perhaps no surprise given its location on the Ghaggar flood plain.

The nearest Himalayan rivers to the Singh et al. (2016) drill sites are the Sutlej and Yamuna that have modern εNd values of −17.0 and −17.7 respectively on the flood plains (Tripathi et al., 2004), with −19 being measured by Alizai et al. (2011) in the Sutlej just before it leaves its headwaters. This compares with an average value of −13.7 for the Thar Desert and −15.6 for the modern Ghaggar. Because the Ghaggar sample of Tripathi et al. (2004) was taken quite far downstream and it is likely that at least some of this value represents recycling of Thar Desert sand rather than the composition of the river as it leaves the headwaters. Although the Yamuna or Sutlej would have been affected by changes in the relative strength of erosion between the Higher and Lesser Himalayas, the composition of the sediment studied here is not consistent with either of these rivers being the source because the εNd values are too positive.

The record of Singh et al. (2016) shows a rise in εNd values from around −17 at 17 ka to around −15 at 5 ka. Such a change in isotopic character could be explained by drainage capture. In this scenario, the pre-17 ka section shows the influence of a west flowing river (Saini et al., 2009), possible the Sutlej or even the Yamuna River, whose capture to the modern eastward flow is loosely dated at some time between 49 and 10 ka (Clift et al., 2012). Thus the trends reconstructed by Singh et al. (2016) may reflect the loss of the Sutlej or Yamuna after 17 ka, together with recycling from the Thar Desert that sharply increases after 10 ka, the latter being driven by the weakening strength of the South Asian monsoon. The trend to more positive εNd values in the Ghaggar through the Holocene does not reflect change in erosion patterns in the mountains themselves, which the delta records show to be the opposite of what is proposed in this study. Alizai et al. (2011) interpreted single grain Pb isotope data from the Hakkra that parallels the Nd isotopes published by East et al. (2015) to reflect evolution of a Himalayan fed stream being clogged by Thar Desert sand as the stream dried up (Giosan et al., 2012). Such a change does not imply a major change of erosion in the source, as proposed by Singh et al. (2016).

If this alternative explanation is accepted then it resolves the apparent discrepancy between the delta and the Ghaggar floodplain isotopic

Fig. 1. Evolution in Nd isotopic character of the whole Indus basin as recorded at the delta, compared to the compositional evolution in the floodplains of the Ghaggar-Hakkra. The results from Singh et al. (2016) are consistent with the earlier dated from the floodplains in modern Pakistan forming a continuous floodplain but from a ephemeral river that did not drain the Lesser or Higher Himalaya. The trend towards desert dune compositions reflects the progressive demise of this river and its choking by the expansion of the Thar Desert during the Holocene.

Fig. 2. Annotated satellite map of NW India and Pakistan showing the location of the modern rivers, together with the drill sites mentioned in the text, as well as the paleo-channels proposed by Saini et al. (2009), Sinha et al. (2013) and Clift et al. (2012).
records. Otherwise, it would be difficult to understand why only the Ghaggar would show more Higher Himalayan erosion during post-glacial times when the Indus Basin on a larger scale shows the reverse. Our alternative explanation would also reconcile the trends in regional erosion with the idea that erosion during dry glacial times was focused in the glaciated Himalaya, while the wetter climate after glaciation shifted the focus into Lesser Himalaya (Clift et al., 2008). In any case, glaciation is generally thought to increase rather than decrease erosion (Hallet et al., 1996) so a trend implying a larger Higher Himalayan flux during deglaciation proposed in the Ghaggar flood plains is enigmatic, although in the short term a local switch to more Higher Himalayan compositions could be driven by recycling from glacial lakes.

Recognition that proximal foreland records do not necessarily reflect erosion in the adjacent mountain chain is an important step in being able to use these sequences to reconstruct long-term environmental and erosional evolution of the mountains and also emphasize the importance of distal records for understanding orogenic scale responses to climate change.

References


