

## Deformation of Rocks: Big Ideas

- Earth scientists use the structure, sequence, and properties of rocks, to reconstruct events in Earth's history
- Understanding geologic processes active in the modern world is crucial to interpreting Earth's past
- Over Earth's vast history, both gradual and catastrophic processes have produced enormous changes
  - Super-continents formed and broke apart
  - mountains formed and eroded away

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## Deformation of Rocks

- Folds and faults are geologic structures caused by deformation.
- Structural geology is the study of the deformation of rocks and its effects.

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Fig. 7.1

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## Orientation of Deformed Rocks

We need some way to describe the distribution of geologic structures. So we use the terms strike and dip.

**Strike:** compass direction of a rock layer as it intersects with a horizontal surface.

**Dip:** acute angle between the rock layer and the horizontal surface, measured perpendicular to strike.

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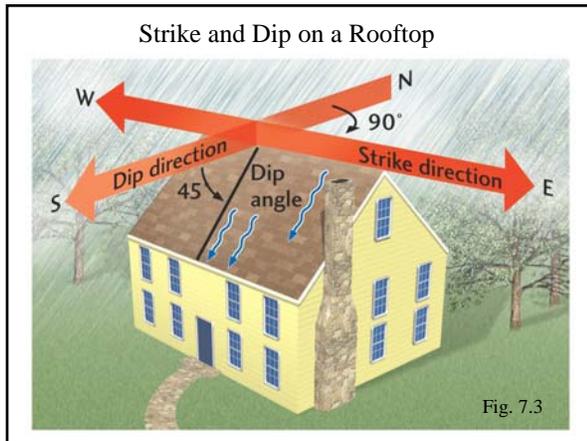
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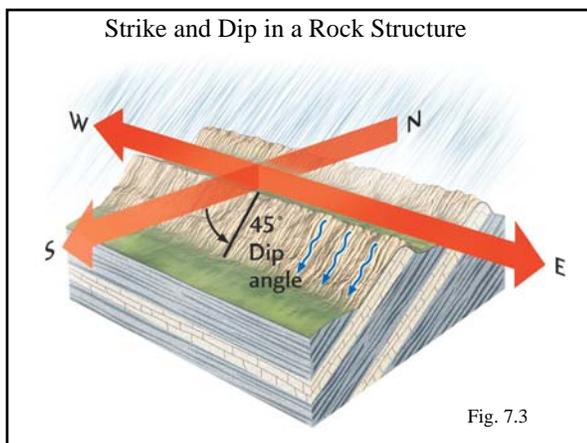
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### Strike and Dip in a Rock Structure



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### Stress: force per unit area

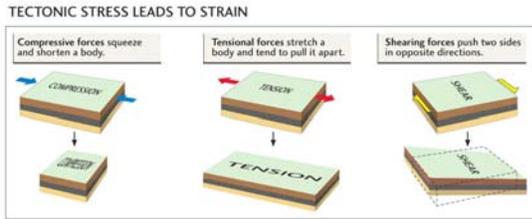


Fig. 7.7

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### Strength

- ability of an object to resist deformation
- compressive or tensile



Fig. 7.5

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## Strain

**Any change in original shape or size of an object in response to stress acting on the object**



Fig. 7.1

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## Elastic Deformation

***Temporary* change in shape or size that is recovered when the deforming force is removed**

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## Ductile (Plastic) Deformation

- **Permanent change in shape or size that is not recovered when the stress is removed**
- **Occurs by the slippage of atoms or small groups of atoms past each other in the deforming material**



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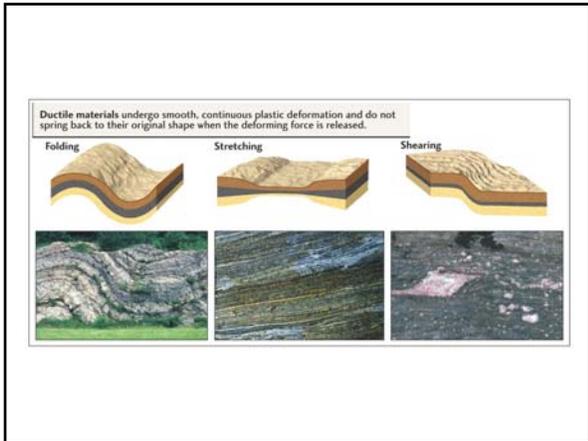
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### Brittle Deformation (Rupture)

- **Loss of cohesion of a body under the influence of deforming stress**
- **Usually occurs along sub-planar surfaces that separate zones of coherent material**

The photograph shows a dark rock face with several prominent, light-colored, sub-planar fracture surfaces (joints or faults) that have broken the rock into blocks.

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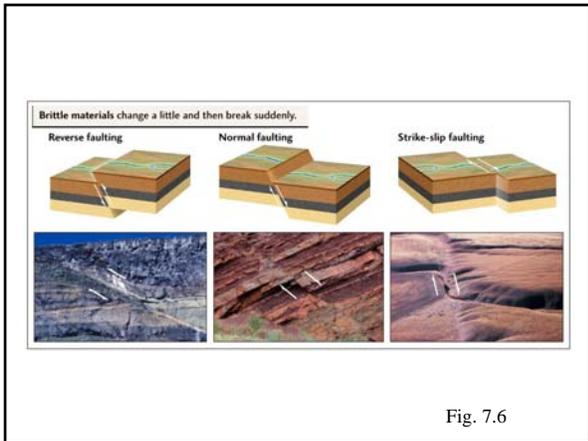


Fig. 7.6

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## Factors that Affect Deformation

- temperature
- pressure
- strain rate
- rock type



*The variation of these factors determines whether a rock will fault or fold.*

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## Joints

**Cracks in rocks along which there has been no appreciable displacement.**



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## Faults

**Fractures in rocks created by earthquakes.**

- Dip-slip faults
  - normal
  - reverse
- Strike-slip faults
- Oblique-slip faults

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Dip-slip Faults: Motion of the fault blocks is parallel to the dip direction.



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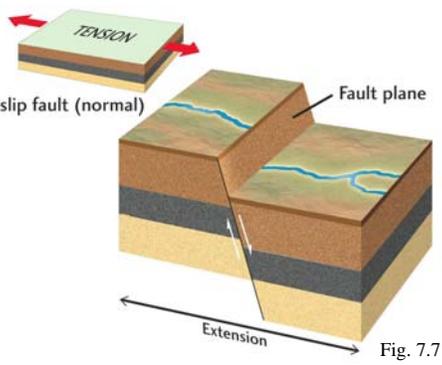
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(a) Dip-slip fault (normal)



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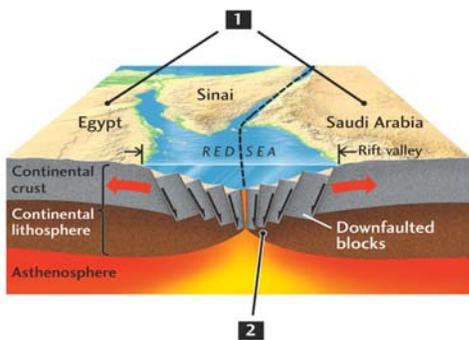
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### Rift Valley



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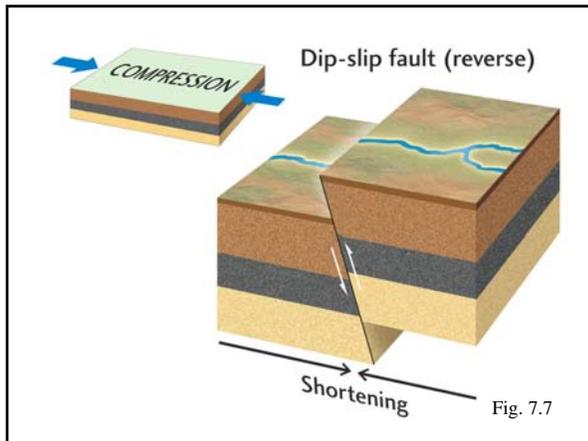
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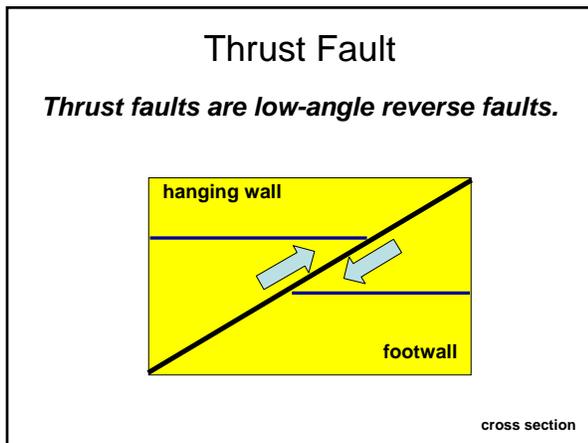
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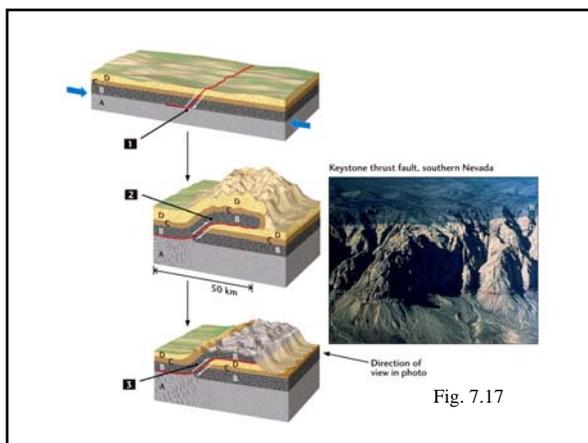
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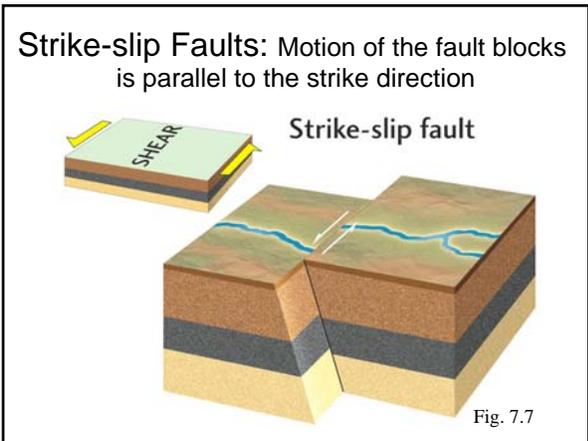
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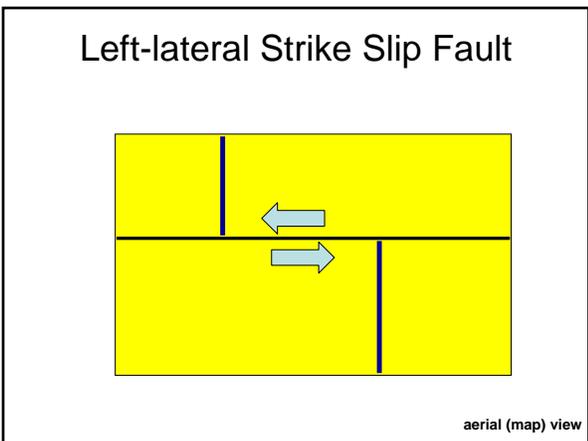
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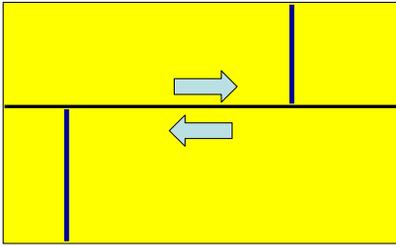
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# Right-lateral Strike Slip Fault



aerial (map) view

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# San Andreas Fault



Fig. 7.6

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# (d) Oblique-slip fault

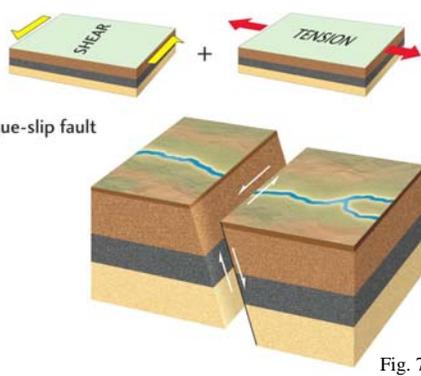


Fig. 7.7

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## Folding of Rocks

- Produced by horizontal or vertical forces
- Scale can be from cm to 100's of km

ROCK FOLDING IS INFLUENCED BY THE TYPE OF ROCK AND THE COMPRESSIVE FORCES

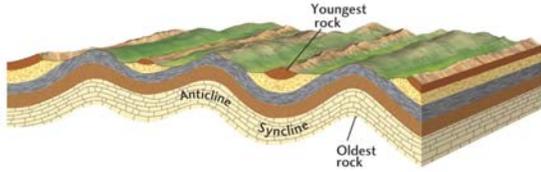


Fig. 7.10

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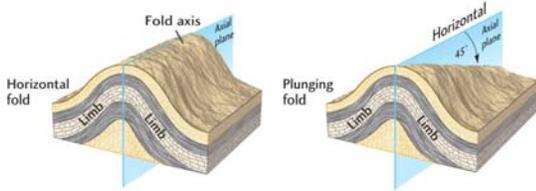
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## Fold Terminology

**axial plane:** the plane of mirror symmetry dividing the fold into two limbs

**axis:** the line formed by the intersection of the axial plane and a bedding plane



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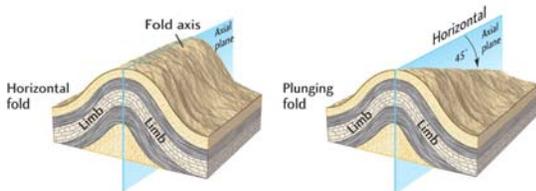
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## Fold Terminology

**horizontal fold:** fold where the axis is horizontal

**plunging fold:** fold where the axis is not horizontal



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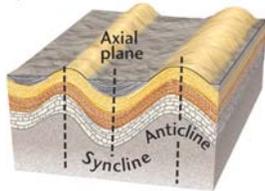
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## More Fold Terminology

**syncline:** a sequence of folded rocks with the youngest rocks on the inside of the fold

**anticline:** a sequence of folded rocks with the oldest rocks on the inside of the fold

Symmetrical folds



Asymmetrical folds

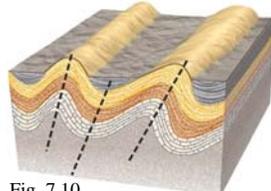


Fig. 7.10

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Fig. 7.10

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## Overturned folds

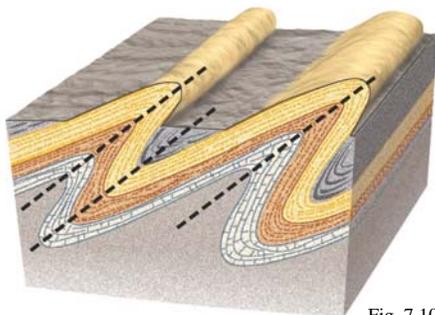


Fig. 7.10

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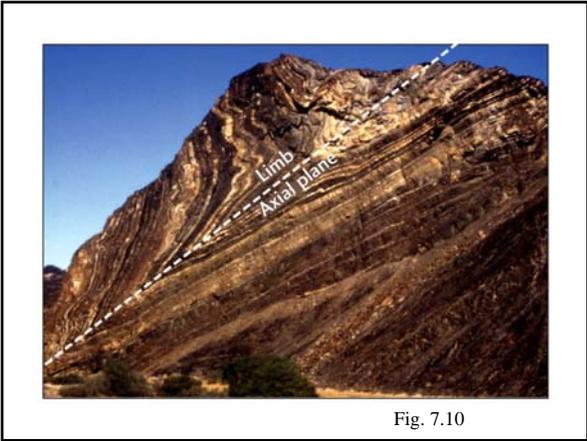


Fig. 7.10

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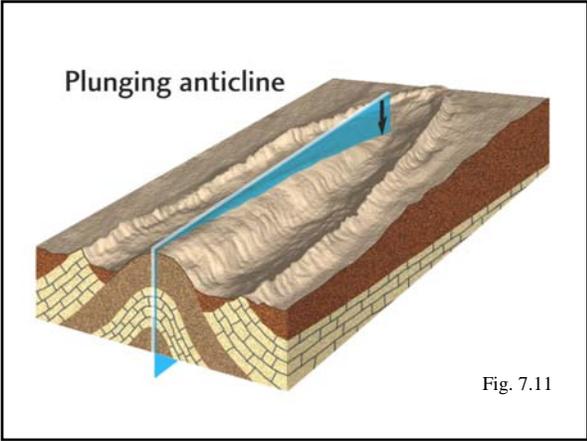


Fig. 7.11

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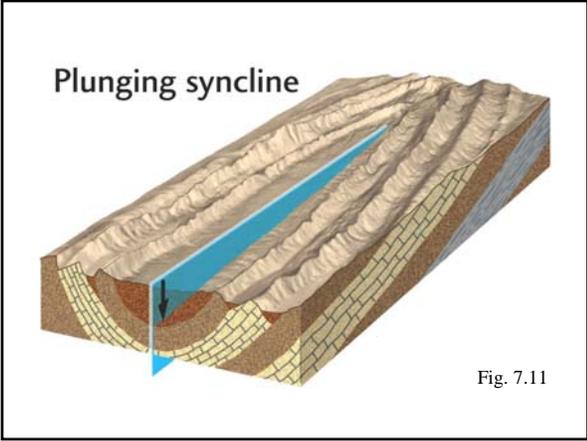


Fig. 7.11

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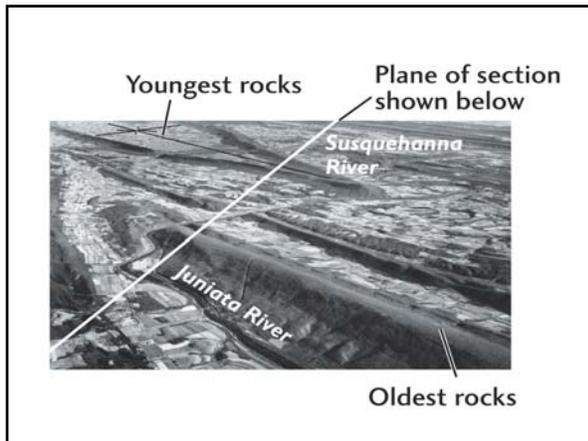
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### And More Fold Terminology

**dome:** a sequence of folded rocks in which all the beds dip away from a central point

**basin:** a sequence of folded rocks in which all the beds dip towards a central point

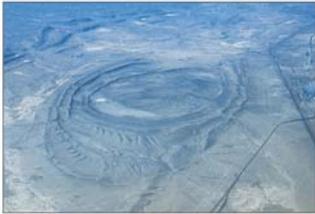


Fig. 7.12

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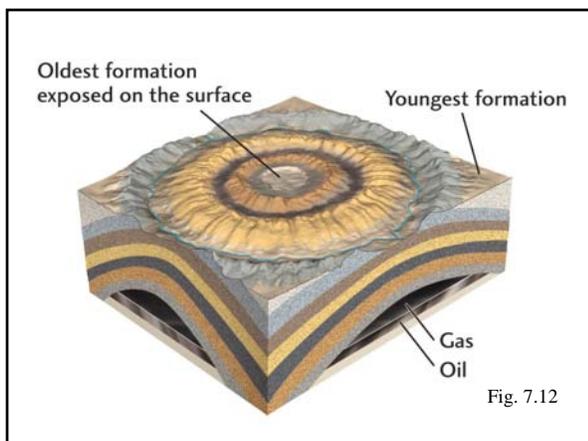
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## Mountain Belts

- narrow zones of folded, compressed rocks with associated magmatism
- formed at convergent plate boundaries
- two major active belts: Cordilleran (Rockies-Andes) and Alps-Himalaya
- older examples include Appalachians and Urals

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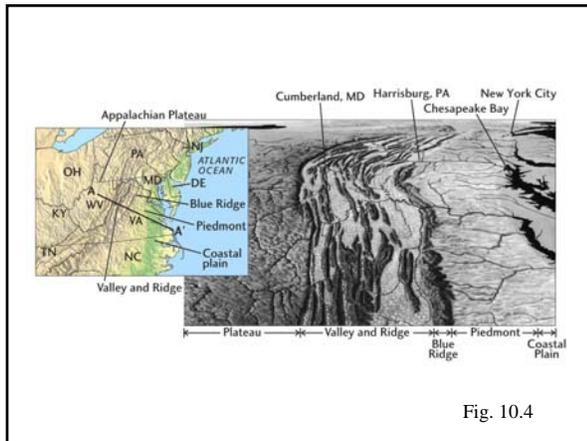


Fig. 10.4

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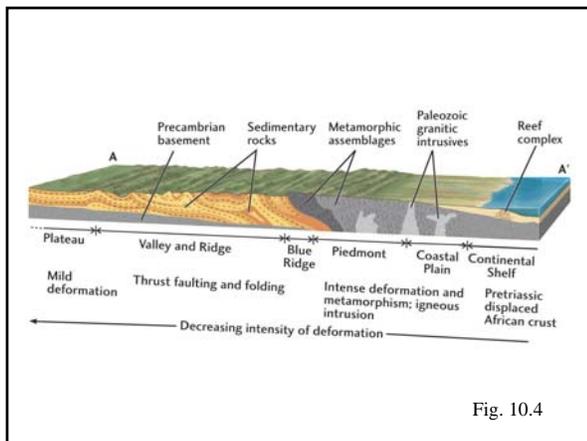


Fig. 10.4

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## North American Cordillera

Complex geologic history from multiple episodes of deformation and magmatism over the past 500 million years.

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## Tectonic Provinces of the West



Fig. 10.5

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## Tectonic History of San Andres Fault

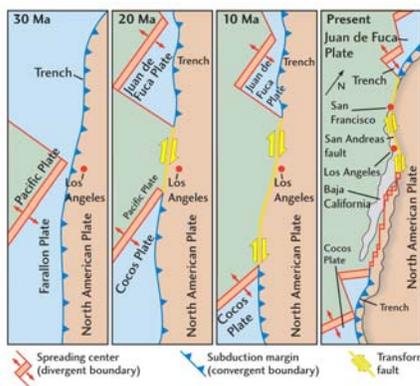


Fig. 10.6

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## Major Uplift Along Normal Faults



Fig. 10.7

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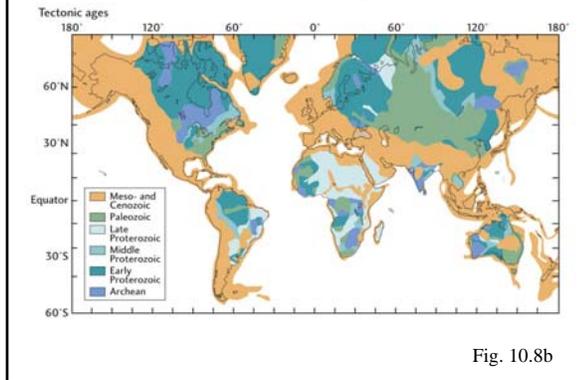
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## Tectonic Ages



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## How Continents Grow

- **Magmatic differentiation:** magma transferred to continents at subduction zones
- **Continental accretion:** buoyant fragments of continents attached to continents as the result of plate motions

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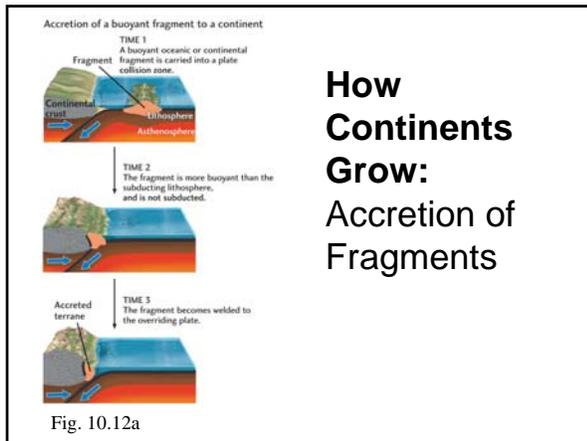
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## How Continents Grow: Accretion of Fragments

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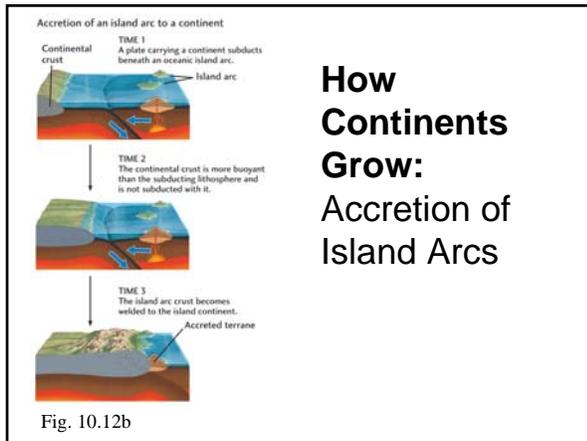
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## How Continents Grow: Accretion of Island Arcs

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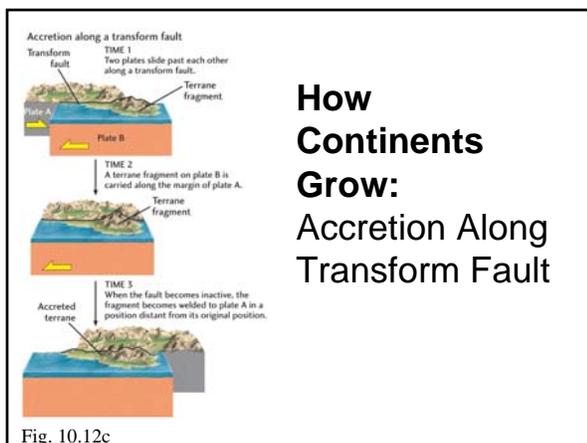
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## How Continents Grow: Accretion Along Transform Fault

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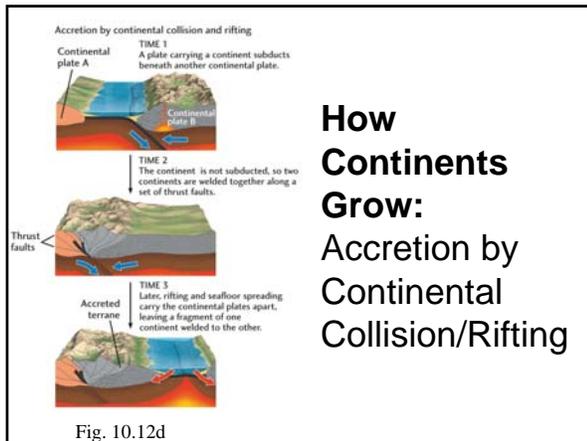
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## How Continents Grow: Accretion by Continental Collision/Rifting

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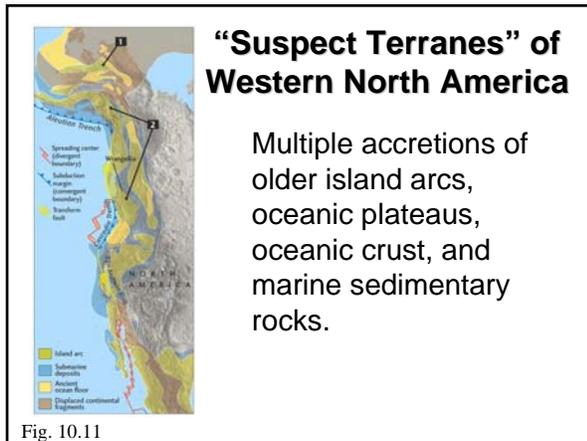
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## “Suspect Terranes” of Western North America

Multiple accretions of older island arcs, oceanic plateaus, oceanic crust, and marine sedimentary rocks.

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## Orogeny

- mountain building
- particularly by folding and thrusting of rock layers
- often accompanied by magmatic activity

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**The Himalayan Orogeny: The Indian Plate subducts under the Eurasian Plate.**

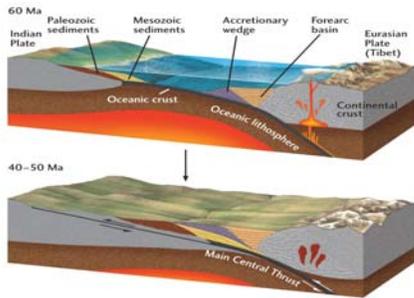


Fig. 10.15

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**The Himalayan Orogeny: India subcontinent collides with Tibet and breaks along the Main Central Thrust fault.**

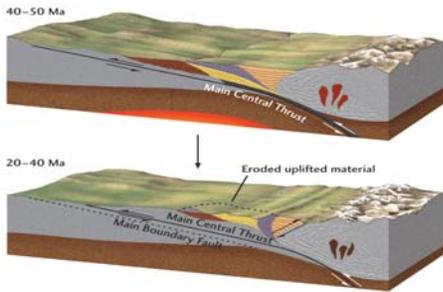


Fig. 10.15

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**The Himalayan Orogeny: A second thrust fault forms, lifting the first fault.**

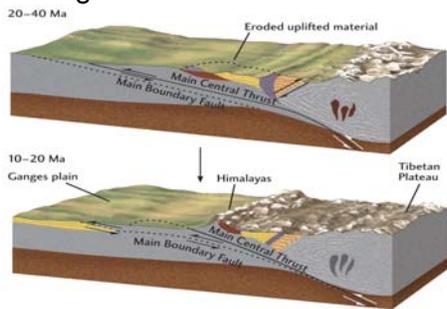


Fig. 10.15

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### Tectonic Features: Collision of India and Eurasia

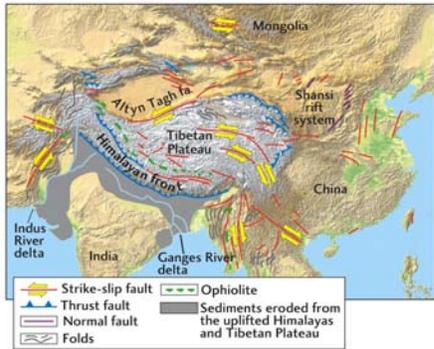


Fig. 10.16

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### The Wilson Cycle

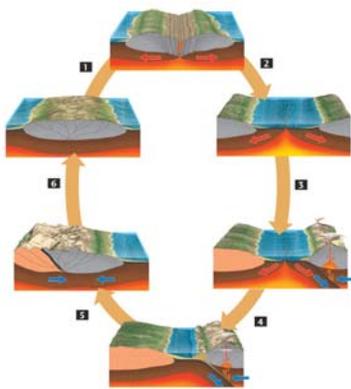


Fig. 10.18

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