

Structure of the Earth

- seismic velocity depends on the composition of the type of material and pressure
- when waves move from one type of material to another, they change speed and direction
- we can use the behavior of seismic waves to tell us about the interior of the Earth



Fig. 14.1

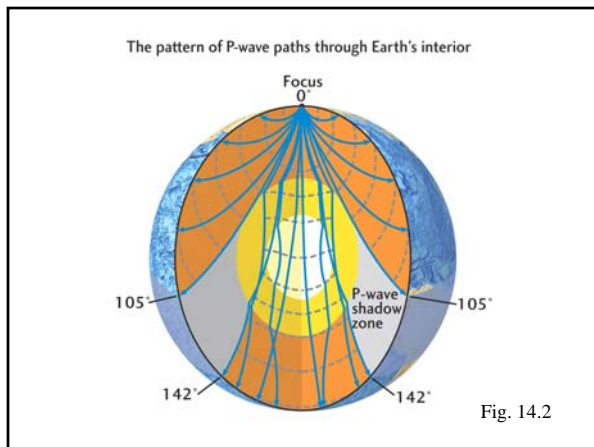
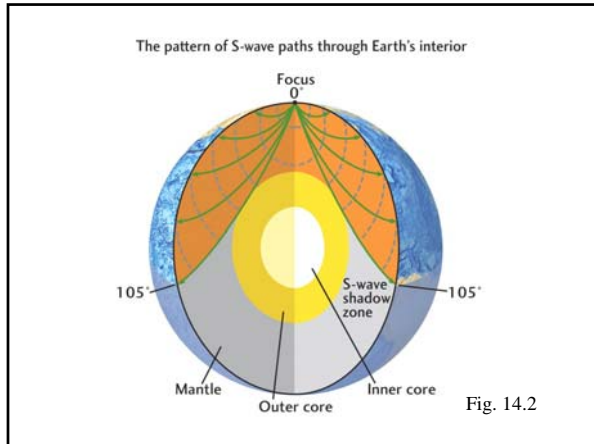
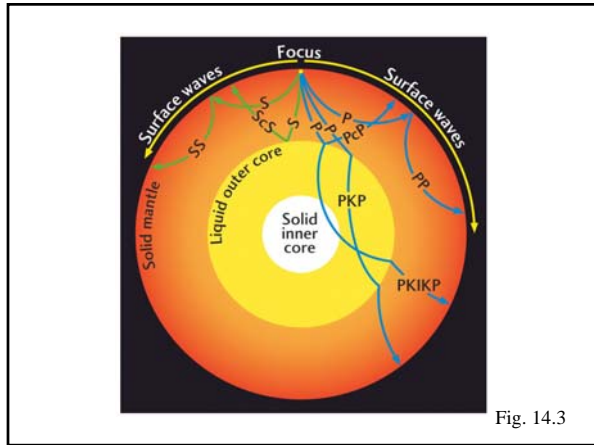
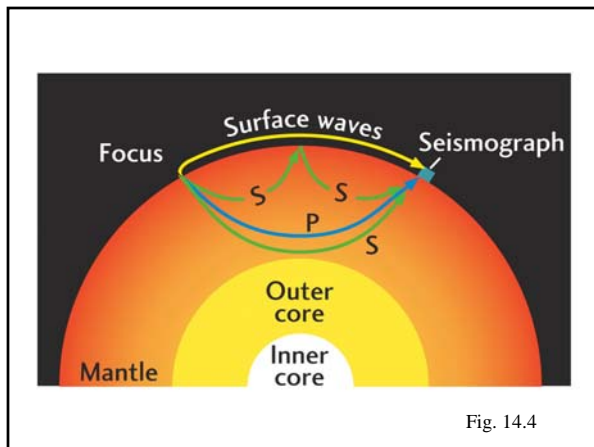


Fig. 14.2







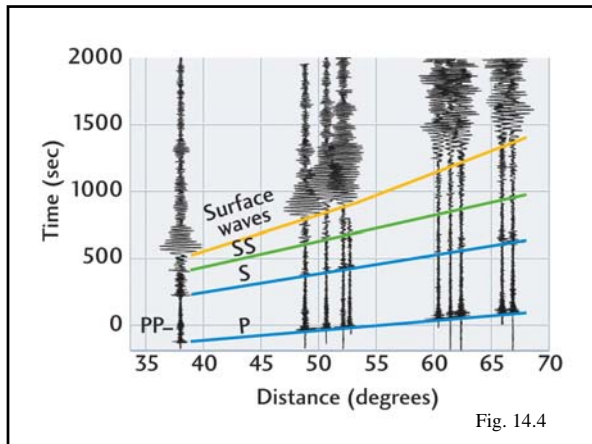


Fig. 14.4

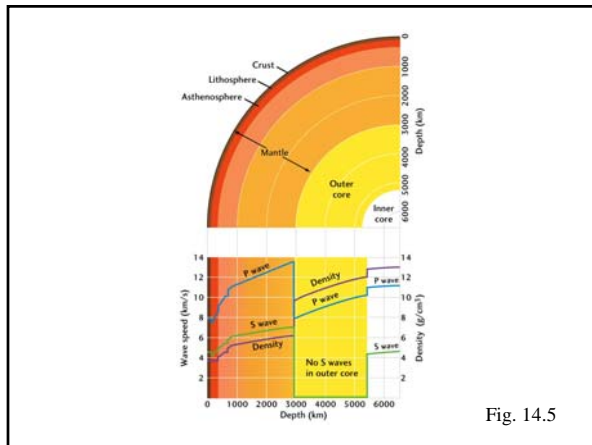


Fig. 14.5

Structure of the Earth

The study of the behavior of seismic waves tells us about the shape and composition of the interior of the Earth:

- *crust*: ~10–70 km, thick, intermediate composition
- *mantle*: ~2800 km, thick, mafic composition
- *outer core*: ~2200 km, thick liquid iron
- *inner core*: ~1500 km, thick solid iron

Composition of the Earth

Seismology also tells us about the density of rocks:

- *continental crust*: $\sim 2.8 \text{ g/cm}^3$
- *oceanic crust*: $\sim 3.2 \text{ g/cm}^3$
- *asthenosphere*: $\sim 3.3 \text{ g/cm}^3$

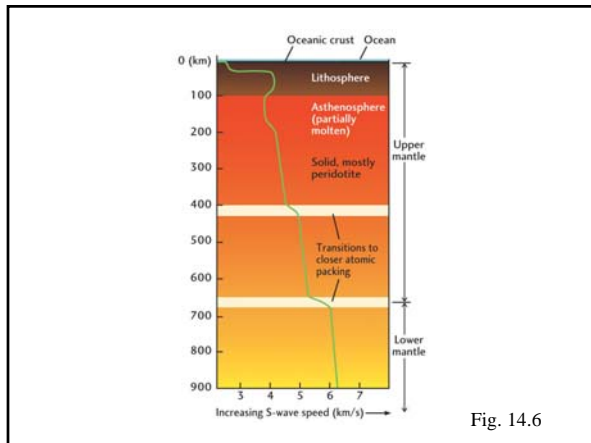


Fig. 14.6

Isostasy

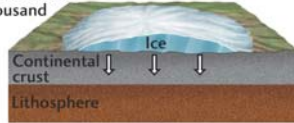
- buoyancy of low-density rock masses “floating on” high-density rocks; accounts for “roots” of mountain belts
- first noted during a survey of India
- Himalayas seemed to affect plumb

Isostasy

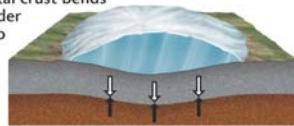


In order for continents to be higher they must also be thicker.

TIME 1
A continental glacier starts to form, and continues to thicken over a few thousand years at the start of an ice age.

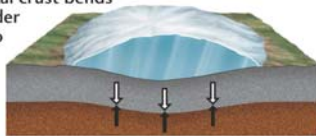


TIME 2
The continental crust bends downward under the ice load to the extent needed to provide buoyant support.

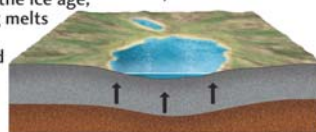


Box 14.1

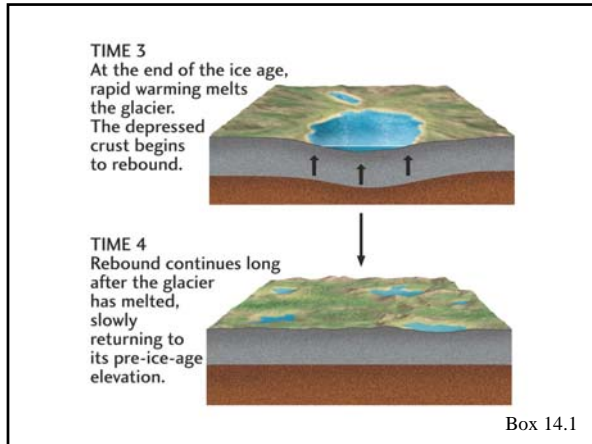
TIME 2
The continental crust bends downward under the ice load to the extent needed to provide buoyant support.



TIME 3
At the end of the ice age, rapid warming melts the glacier. The depressed crust begins to rebound.

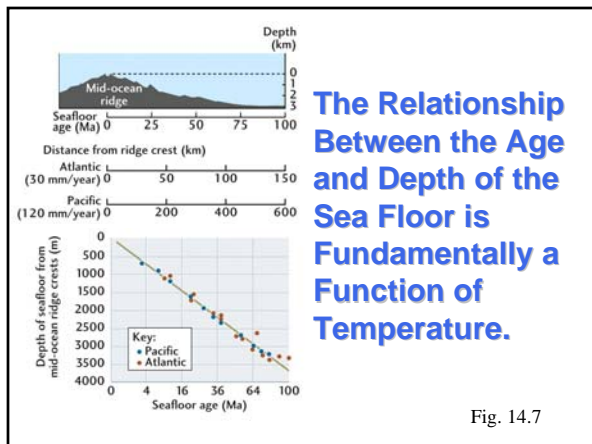


Box 14.1



Earth's Internal Heat

- original heat
- subsequent radioactive decay
- conduction
- convection



Temperature Increases With Depth: The Geotherm Curve

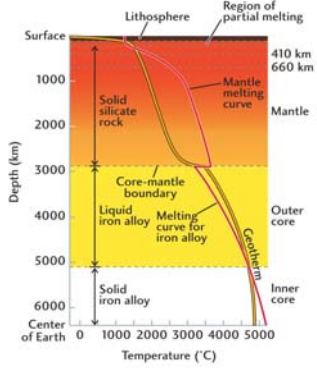


Fig. 14.8

Seismic Tomography Uses Travel Times to Create 3-D Images of Earth's Interior

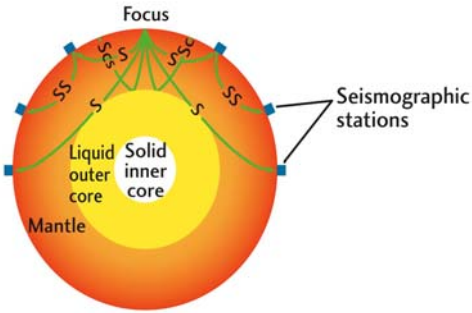


Fig. 14.9

Tomographic Section Reveals Hot and Cold Rocks

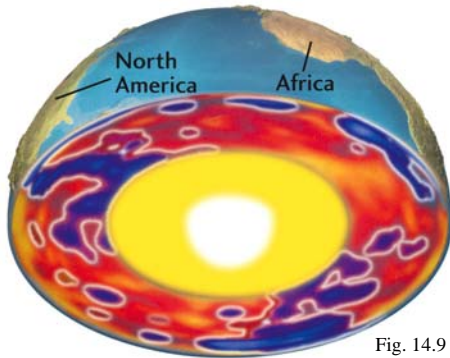
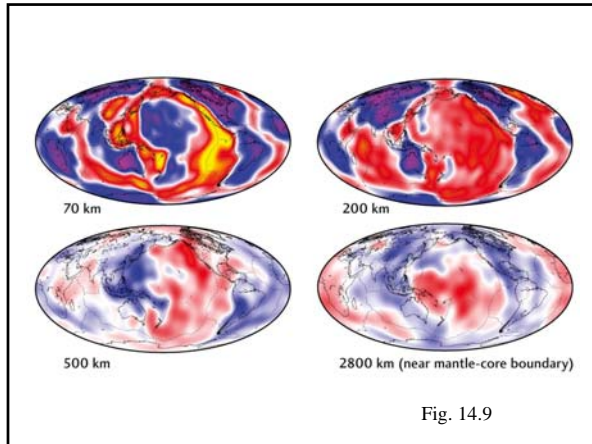
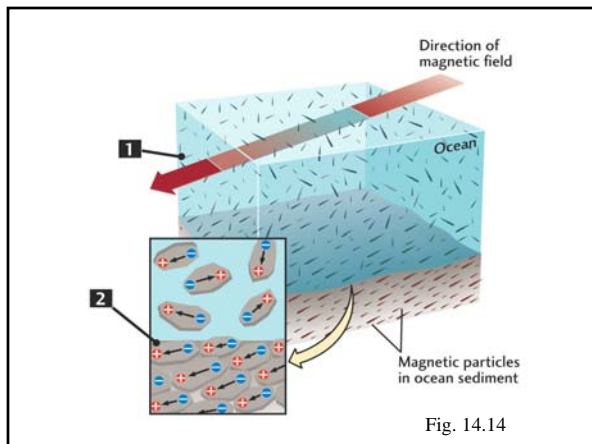


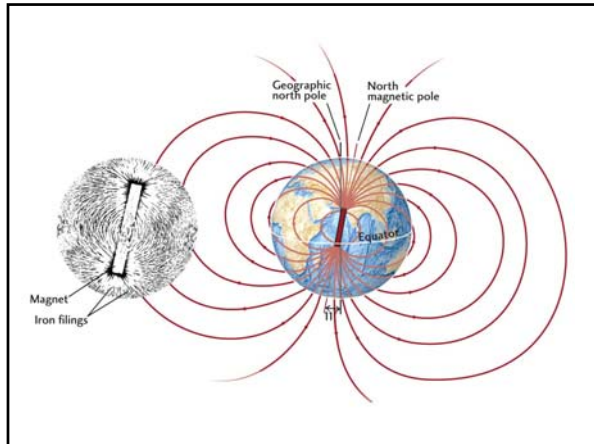
Fig. 14.9



Paleomagnetism

- use of the Earth's magnetic field to investigate past plate motions
- permanent record of the direction of the Earth's magnetic field at the time the rock was formed
- may not be the same as the present magnetic field

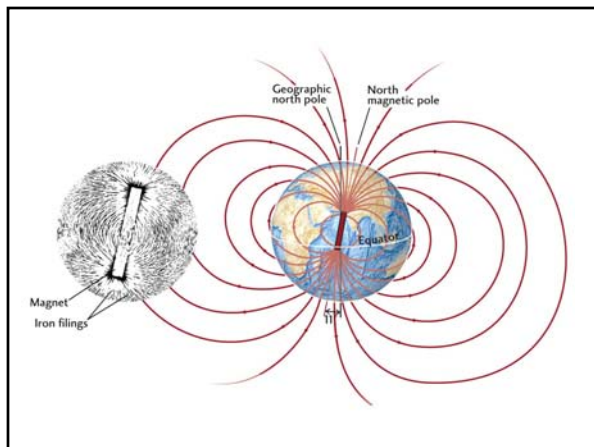




Earth's Magnetic Field

declination: horizontal angle between magnetic north and true north

inclination: angle made with horizontal



Earth's Magnetic Field

- It was first thought that the Earth's magnetic field was caused by a large, permanently magnetized material deep in the Earth's interior.
- In 1900, Pierre Currie recognized that permanent magnetism is lost from magnetizable materials at temperatures from 500 to 700 °C (Currie point).

Geodynamo: Self-exciting Dynamo

A dynamo produces electric current by moving a conductor in a magnetic field and *vice versa*. (*i.e.*, an electric current in a conductor produces a magnetic field.

Geodynamo: Self-exciting Dynamo

- It is believed that the outer core is in convective motion (because it is liquid and in a temperature gradient).
- A "stray" magnetic field (probably from the Sun) interacts with the moving iron in the core to produce an electric current that is moving about the Earth's spin axis yielding a magnetic field—a self-exciting dynamo!

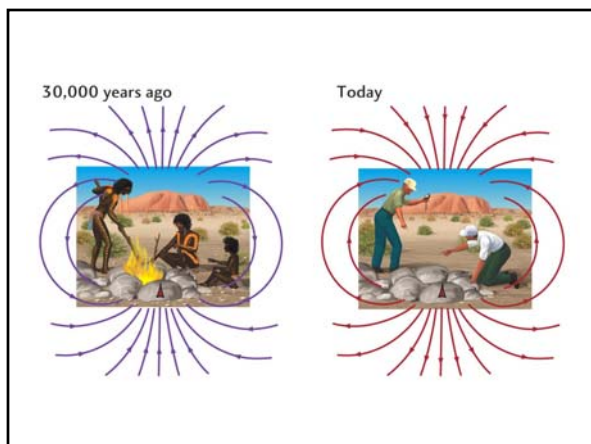
Geodynamo: Self-exciting Dynamo

The theory is gaining popularity because:

- it is plausible
- it predicts that the magnetic and geographic poles should be nearly coincident
- the polarity is arbitrary
- the magnetic poles move slowly

Magnetic Reversals

- the polarity of the Earth's magnetic field has changed thousands of times in the Phanerozoic
- The most recent reversal was about 30,000 years ago. However, the end of the last significant reversal was approximately 700,000 years ago.
- These reversals appear to be abrupt, lasting approximately 1000 years.



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Magnetism and Dating

We can now use the magnetic properties of a sequence of rocks to determine their age.

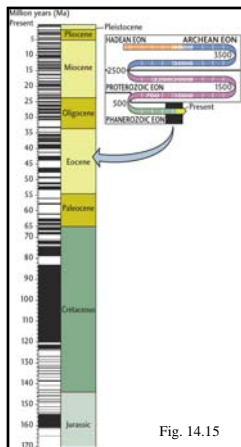


Fig. 14.15

The Geomagnetic Time Scale
Based on determining the magnetic characteristics of rocks of known age (from both the oceans and the continents).

We have a good record of geomagnetic reversals back to about 60 Ma.
